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Total, cold and hot water extractable organic carbon in soil profile: impact of land-use change

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Abstract

The content of labile, especially water extractable organic carbon (WEOC) is a sensible indicator of soil organic matter quality. The main objectives of this study were: i) to investigate the profile changes of cold and hot water extractable organic carbon in forest and arable soils; ii) to evaluate the correlation between these labile fractions of soil organic matter and total organic carbon content. The experiments were carried out on a *Gleyic Albeluvisol* (ABg) in the upper part of Dniester basin, Western Ukraine. The soil samples were taken from 50-cm depth soil profile with 5-cm step. Total organic carbon (TOC), cold water extractable organic carbon (CWEOC) and hot water extractable organic carbon (HWEOC) contents in soil were determined as well as pH (H₂O) and electrical conductivity of soil:water suspensions. The results of this study showed that in 0–50 cm layer of arable soil TOC content decreased by 32%, CWEOC – by 23% and HWEOC – by 74% compared to forest soil that confirmed a high informative role of HWEOC fraction. The profile changes of WEOC percentage were analysed. They also show that HWEOC is much more informative indicator of soil organic matter quality than CWEOC. The most prominent changes of soil chemical properties, TOC, CWEOC and HWEOC contents in response to deforestation were observed in the top 5-cm soil layer. We suggested this thin soil layer to be defined as soil stress-sensitive zone.

Key words: dissolved organic carbon, labile fraction, soil organic matter, total organic carbon, water-soluble organic matter.

Introduction

Soil carbon reservoir is one of the largest on the planet and contains up to 2500 Pg (petagram, 1 Pg = 10¹⁵ g) C in upper 1 m layer (Batjes, 1996). It is three times larger than atmospheric carbon reservoir (Oelkers, Cole, 2008) and contains almost 4.5 times the amount of carbon in biomass (Lal, 2008). Thus the quality and quantity of soil organic matter (SOM) play the key role in global carbon balance. The most labile fractions of SOM are water-soluble organic substances (Bu et al., 2011). Due to high biodegradation rates they are an important source of carbon dioxide production and directly affect soil CO₂ efflux (Gregorich et al., 2003; Kim et al., 2012).

Water extractable organic carbon (WEOC) is an important but not homogenous pool. Solubility of different organic compounds depends on the temperature of solvent. Therefore two functional pools of WEOC are usually studied: cold water extractable organic carbon (CWEOC) and hot water extractable organic carbon (HWEOC) (Ghani et al., 2003). The first one quantitatively is very close to dissolved organic carbon measured directly in soil using different lysimeters and suction cups (Rees, Parker, 2005; Ostrowska et al., 2010).

The second one consists of more stable components that form the close reserve of nutrients and energy for plants and microorganisms (Bu et al., 2011). Landgraf et al. (2006) conducted a detailed study on cold and hot water extractable organic matter as indicators of litter decomposition in forest soils. They considered HWEOC to be a heterogeneous pool, because hot (≥70°C) water kills vegetative cells of microorganisms and extracts many components from microbial biomass, as well as many nonmicrobial organic substances. Therefore, HWEOC has much higher biodegradability rates than CWEOC (Gregorich et al., 2003). In the majority of WEOC studies only the uppermost soil layer was analyzed (Shpakivska, 2008; Ahn et al., 2009; Šlepetienė et al., 2010; Ruqin et al., 2013). In fewer cases the profile changes of WEOC were investigated (Boyer, Groffman, 1996; Menyailo, 2008; Nakanishi et al., 2012).

The objectives of this study were: i) to investigate the profile changes of cold and hot water extractable organic carbon in forest and arable soils; ii) to evaluate the correlation between CWEOC, HWEOC and total organic carbon (TOC) in these soils.

Materials and methods

Study area. The experiments were carried out on two plots (100 × 100 m) in the upper part of Dniester basin, Western Ukraine (49°32' N, 23°20' E). The field experimental plots were established in 2006. The first one was set up in a pristine old-growth hornbeam-oak forest. There were English oak (*Quercus robur* L.) – 35%, sycamore maple (*Acer pseudoplatanus* L.) – 1% in tree stratum on the first floor and European hornbeam (*Carpinus betulus* L.) – 25%, small-leaved lime (*Tilia cordata* Mill.) – 3%, field maple (*Acer campestre* L.) – 1%, mountain-ash (*Sorbus aucuparia* L.) – 1% in the second one. The shrub layer was represented by common hazel (*Corylus avellana* L.) – 7% and common dogwood (*Swida sanguinea* L.) – 5%; the herb layer – by more than 50 species among which yellow archangel (*Galeobdolon luteum* Huds.) – 15%, wood sorrel (*Oxalis acetosella* L.) – 10% and European woodland sedge (*Carex sylvatica* Huds.) – 5% were the most common. The second plot was set up in an adjacent tillable land, where the forest was clear cut about 70 years ago. The four-field crop rotation was applied: potatoes, barley + clover, clover, winter wheat. The soil (*Gleyic Albeluvisol*, ABg) was fertilized with cattle manure (40 t ha⁻¹ once in four years) and mineral fertilizers (N₄₅P₄₅K₅₀). For the last 30 years mean annual total precipitation was 697 mm; the annual average temperature was 7.8°C, and the sum of active temperatures – about 2400–2600°C.

Soil sampling and analysis. Five soil pits were made on each experimental plot. For this study the soil samples were taken from 50-cm depth soil profile, with a 5-cm step. Soil sampling was done in October 2008. The samples were taken from three sides of each soil pit and with a special bore from five points located less than 5 m from the pit. Fresh soil samples were passed through 3 mm sieve and composite samples for each 5-cm layer were taken. For all the analyses, described in this study, the soil samples were air-dried. The chemical composition was provided in dry matter. Total organic carbon (TOC) content was determined by wet combustion method (ISO 14235: 1998, Soil quality – Determination of organic carbon by sulfochromic oxidation). The absorbance of the obtained solutions was measured spectrophotometrically on a double beam UV-visible spectrophotometer SPEKOL 2000 (“Analytik Jena”, UK). Content of water extractable organic carbon (WEOC) was determined according to the method of Haynes and Francis (1993) modified by Ghani et al. (2003) that consists of two-step water extraction. A 3 g sample of air dried soil was put into 50 ml polypropylene centrifuge tube and 30 ml of distilled water were added. The tube was put on end-over-end shaker at 30 rpm for 30 min at 20°C. After extraction, the suspension was centrifuged at 3500 rpm for 20 min and the supernatant solution was filtered through 0.45 µm membrane filters. A 5 ml portion of the obtained solution was put into a glass tube and evaporated to dryness at 60°C. The fraction obtained is cold water extractable organic carbon (CWEOC). A 20 ml portion of distilled water was added to the same centrifuge tube with sediments. The used membrane filter was washed with 10 ml of distilled water, which was collected into the same tube. It was capped, shaken vigorously for 10 s and put into the thermostat at 80°C for 16 h. After the extraction, the tube was shaken for 10 s once again. The

suspension was centrifuged, filtered and the solution was evaporated like in previous step. Organic carbon content was determined by wet combustion method as described before. In each experiment, five replicates were taken for each soil sample. Electrical conductivity and pH (potentiometric) were measured in 1:2.5 soil:water suspensions on a conductivity meter “S47 SevenMulti” (“Mettler-Toledo”, USA). Soil particle size distribution was measured with sedimentation technique of Kaczynski (1965).

Statistical analysis. Statistical analysis of experimental data was made according to the recommendations of Glantz (1997). For comparisons of multiple groups the Conover post hoc test (Conover, 1999) was used. The difference was considered significant when $P < 0.05$. Linear and exponential regressions analysis was used to model the profile changes of TOC, CWEOC and hot water extractable organic carbon (HWEOC) contents in soil with depth. The linear regression was used to model the relationship between TOC, CWEOC and HWEOC contents in soil. Pearson correlation coefficient was used as a measure of the strength of linear dependence between studied parameters (Rodgers, Nicewander, 1988). Statistical analyses were performed using the *MS Excel 2010* with add-in *AtteStat 12.5*.

Results and discussion

All of the measured basic chemical properties of studied forest soil displayed a dependency on soil depth (Table 1). The average TOC content in the forest soil was 15.6 mg g⁻¹ C in the studied 0–50 cm layer. A significant decrease of TOC content with depth was observed from 59.78 in top 5-cm soil layer to 2.53 mg g⁻¹ C at the depth 45–50 cm. This dependence can be approximated by the exponential equation $y = 43.914e^{-0.06x}$ ($P < 0.01$; $n = 10$). Here and elsewhere we used medians (5 soil pits) to perform exponential or linear regression analysis. The decrease of soil acidity with depth was also observed. The lowest pH (H₂O) values (4.24–4.33) were determined in the top 0–5 cm soil layer. In the lower portion of the profile, at the depth 40–50 cm they reach almost 5.0. The highest values of electrical conductivity (EC) were also confined to the uppermost soil layers, where they reached 75 µSm cm⁻¹. In particular, in the top horizons the electrical conductivity values were approximately twice as high as those in the deeper layers (34 µSm cm⁻¹). TOC content and EC values slightly increased in soil layer 40–45 cm where the zones of illuviation were formed.

Over 70 years of agricultural exploitation of the deforested soil, its properties dramatically changed (Table 2). Well stratified organoprofile of forest soil was simplified. The average TOC content in the forest soil of 0–50 cm layer was 11.9 mg g⁻¹ C. The dependence of TOC content and soil depth can be approximated not by the exponential equation as in previous variant, but by a simple linear function $y = 19.527 - 0.305x$ ($r = -0.97$, $P < 0.001$; $n = 10$). Very similar equation $y = 18.047 - 0.289x$ ($r = -0.9743$, $P < 0.001$; $n = 10$) can be used to model the profile changes of EC in this soil. The absolute values of TOC and EC decreased twice in comparison to forest soils. Regarding pH of studied arable soil, it is more than 10 times greater than in forest soil: 5.67 versus 4.31 in top 5-cm layer. It slightly increased with depth up to 5.82 in the layer 45–50 cm.

Table 1. Basic chemical properties of forest soil

Depth cm	Total organic carbon (TOC) mg g ⁻¹ C			pH (H ₂ O)			Electrical conductivity (EC) μSm cm ⁻¹		
	med	max	min	med	max	min	med	max	min
0–5	59.78	69.50	51.00	4.31	4.33	4.24	75.0	63.0	83.0
5–10	26.70	29.61	21.93	4.35	4.39	4.30	40.0	38.0	46.0
10–15	17.69	19.24	14.82	4.40	4.47	4.39	42.0	40.0	44.0
15–20	12.86	14.62	10.54	4.47	4.55	4.45	27.0	24.0	31.0
20–25	11.87	13.42	10.02	4.59	4.62	4.54	37.0	34.0	41.0
25–30	10.22	12.29	8.92	4.61	4.62	4.55	34.0	32.0	39.0
30–35	7.14	8.16	6.51	4.67	4.71	4.58	35.0	34.0	40.0
35–40	3.08	4.03	2.41	4.80	4.82	4.74	27.0	23.0	29.0
40–45	4.18	5.16	3.15	4.95	5.04	4.93	45.0	41.0	48.0
45–50	2.53	3.03	1.90	4.92	4.97	4.86	34.0	32.0	39.0

med – median, max – maximum, min – minimum; n = 5 for each soil layer

Table 2. Basic chemical properties of arable soil

Depth cm	Total organic carbon (TOC) mg g ⁻¹ C			pH (H ₂ O)			Electrical conductivity (EC) μSm cm ⁻¹		
	med	max	min	med	max	min	med	max	min
0–5	19.48	20.55	18.30	5.67	5.71	5.60	34.0	32.5	38.0
5–10	16.45	18.04	16.10	5.55	5.60	5.48	30.0	26.0	32.0
10–15	17.29	18.58	15.15	5.60	5.61	5.54	29.0	26.0	32.0
15–20	14.85	15.50	13.05	5.62	5.71	5.53	27.0	23.0	29.0
20–25	10.85	11.51	9.82	5.64	5.73	5.61	25.0	21.0	26.0
25–30	10.05	10.23	8.55	5.70	5.78	5.69	26.0	25.0	29.0
30–35	9.05	9.55	8.25	5.73	5.79	5.70	30.0	27.0	32.5
35–40	7.80	8.55	7.14	5.81	5.85	5.77	31.0	30.0	35.0
40–45	7.18	7.60	6.23	5.78	5.84	5.71	28.0	24.0	31.5
45–50	6.03	6.90	5.70	5.82	5.91	5.80	27.5	24.0	32.0

med – median, max – maximum, min – minimum; n = 5 for each soil layer

All the basic chemical properties are strongly related to particle size distribution. It affects the reactivity of solids in soil which participate in various chemical reactions. Soil texture may influence the acidity values in soil profile and the rates of accumulation

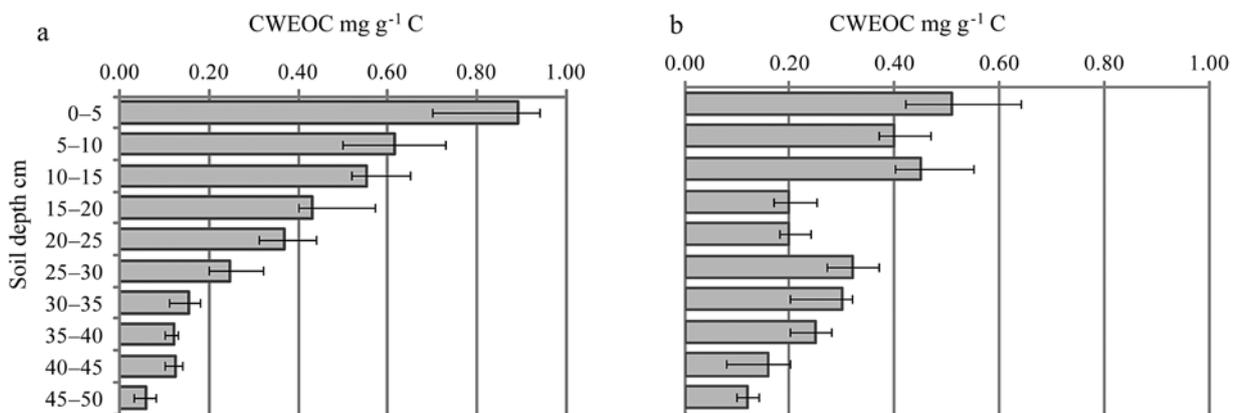
and mineralization of soil organic matter (SOM). Clay content in soil may provide important information about the pattern of illuviation. Comparison of data in Tables 1, 2 and 3 showed that pH (H₂O), TOC and clay content often change in studied soil profiles simultaneously.

Table 3. Particle size distribution of forest and arable soils

Depth cm	Forest soil			Arable soil		
	sand % (1.0–0.01 mm)	silt % (0.01–0.001 mm)	clay % (<0.001 mm)	sand % (1.0–0.01 mm)	silt % (0.01–0.001 mm)	clay % (<0.001 mm)
0–5	23.0	71.4	5.6	39.2	52.1	8.7
5–10	26.8	63.5	9.7	36.9	56.0	7.1
10–15	27.9	64.3	7.8	37.9	55.2	6.9
15–20	29.4	62.5	8.1	39.1	54.1	6.8
20–25	30.9	53.1	16.0	50.4	42.4	7.3
25–30	29.1	61.9	9.0	58.4	25.0	16.6
30–35	30.0	58.9	11.0	51.3	30.9	17.8
35–40	30.8	58.0	11.2	41.7	44.0	14.3
40–45	36.0	54.2	9.8	53.4	26.7	19.9
45–50	29.4	61.5	9.1	64.9	15.9	19.2

Small sizes of soil particles cause their high sorptive potential. Therefore, the zones of high clay content should be rich in water extractable organic matter. In forest soil the average CWEOC content in 0–50 cm layer was $0.36 \text{ mg g}^{-1} \text{ C}$. The highest amount of cold water extractable organic matter (Fig. 1) was located in the uppermost layer 0–5 cm ($0.7\text{--}0.89 \text{ mg g}^{-1} \text{ C}$). CWEOC content gradually decreased with depth to $0.03\text{--}0.08 \text{ mg g}^{-1} \text{ C}$ in the layer 45–50 cm. This dependence can be approximated by the exponential equation $y = 1.0856e^{-0.057x}$ ($P < 0.001$; $n = 10$). In contrast, profile changes of CWEOC in arable soil had different pattern. In comparison to forest soil, it can be seen that deforestation and agricultural exploitation caused significant decrease of CWEOC in top soil layer to 15 cm depth. The average content of CWEOC in 0–50 cm soil layer decreased to $0.29 \text{ mg g}^{-1} \text{ C}$. Another important observation is that a zone of CWEOC accumulation was formed at the depth of 25–40 cm. In this layer clay content is the highest (16.6–19.9%) in studied 50-cm soil profile.

A possible explanation for this might be that ploughing depth was about 25 cm. Therefore, after the tillage the most labile fraction of SOM might have been eluted down the soil profile and sorbed on the surface of clay particles (Corvasce et al., 2006). These changes are associated with significant decrease of TOC content in arable soil compared to forest soil. In our previous studies we investigated the impacts of deforestation on readily oxidizable organic carbon content in forest and arable soils (Partyka, Hamkalo, 2010). Here we found that the profile changes of CWEOC and the content of organic carbon hydrolyzed with 10% H_2SO_4 have much in common – content of both these fractions decrease dramatically in uppermost layer of deforested soil. These findings further support the idea that even principally different SOM extraction techniques may evaluate chemically similar fractions of SOM that form its active pool (von Lützow et al., 2007).

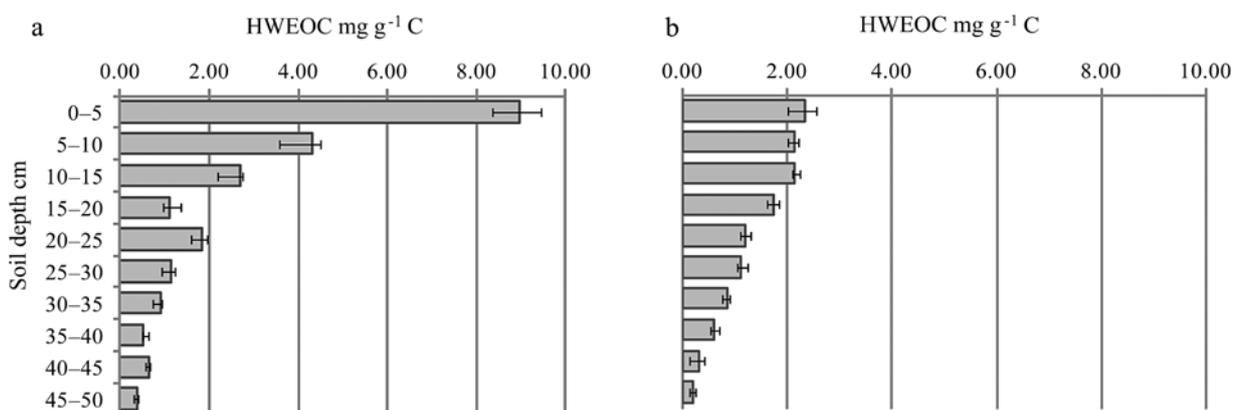


Note. Bars represent median, maximum and minimum; $n = 5$ for each soil layer.

Figure 1. Cold (20°C) water extractable organic carbon (CWEOC) in forest (a) and arable (b) soils

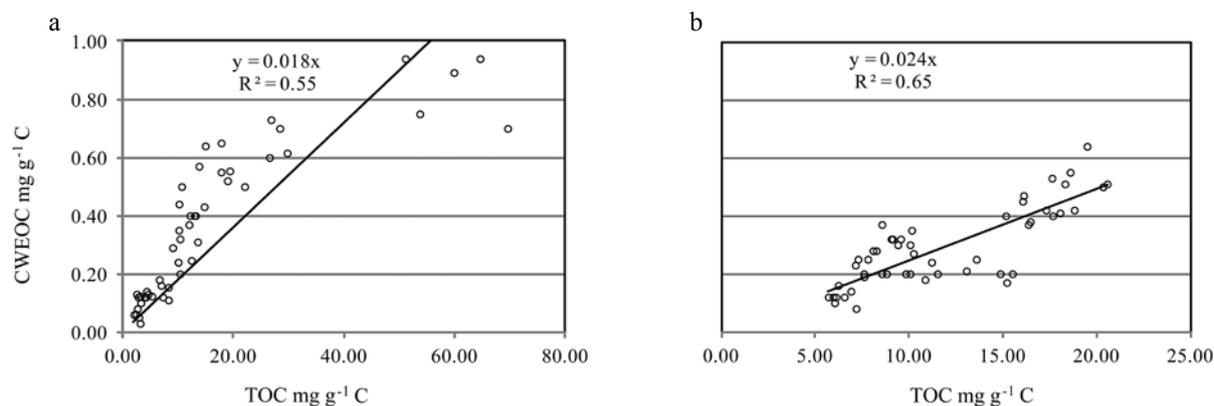
Figure 2 illustrates the profile changes of HWEOC. This data evidently displays that highest amounts ($8.35\text{--}9.43 \text{ mg g}^{-1} \text{ C}$) of this fraction were located in the top 5-cm layer of forest soil. It was 2.1 times higher than at the 5–10 cm depth. HWEOC content gradually decreased with depth to $0.95\text{--}1.34 \text{ mg g}^{-1} \text{ C}$ in 15–20 cm soil layer and to $0.29\text{--}0.41 \text{ mg g}^{-1} \text{ C}$ at the

45–50 cm depth. The dependence of HWEOC content and soil depth can be approximated by the exponential equation $y = 6.5482e^{-0.062x}$ ($P = 0.006$; $n = 10$). Conversely, profile changes of HWEOC in studied arable soil had different pattern. It was only $2.01\text{--}2.55 \text{ mg g}^{-1} \text{ C}$ in top 5-cm soil layer – 3.8 times lower than in forest soil. The dependence between HWEOC content and soil depth can



Note. Bars represent median, maximum and minimum; $n = 5$ for each soil layer.

Figure 2. Hot (80°C) water extractable organic carbon (HWEOC) in forest (a) and arable (b) soils



Note. R^2 – coefficient of determination; both regression lines are statistically significant at the $P < 0.01$ level; $n = 50$.

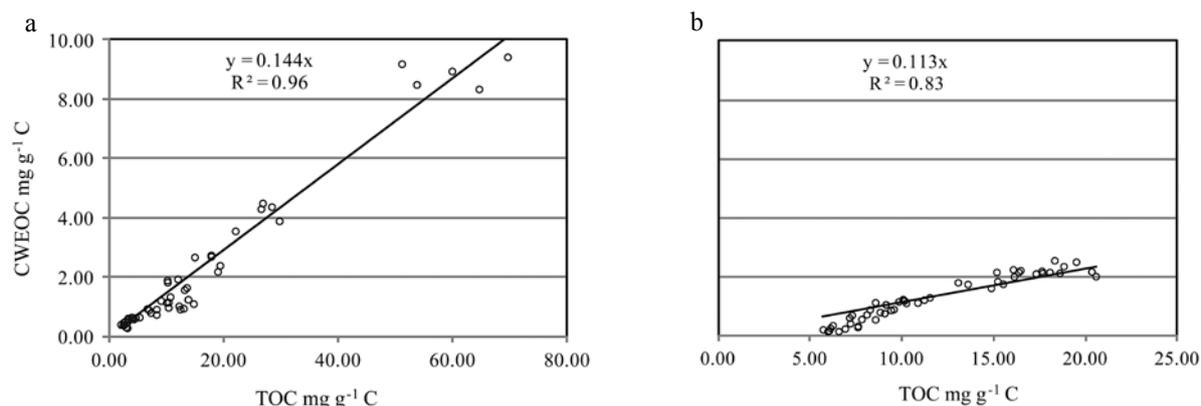
Figure 3. The dependence of cold (20°C) water extractable (CWEOC) and total organic carbon (TOC) contents in forest (a) and arable (b) soils

be approximated by a simple linear function $y = 2.5627 - 0.0517x$ ($r = -0.99$, $P < 0.001$; $n = 10$). The average content of this fraction was 2.24 and 1.27 $\text{mg g}^{-1} \text{C}$ in 0–50 layer of forest and arable soil, respectively. The findings of the current study are consistent with those of Ruqin et al. (2013) who found that HWEOC content in 0–30 cm layer was 58.8% lower in cultivated than in uncultivated soils.

The most prominent changes of basic chemical properties (Tables 1–2), CWEOC and WEOC contents in studied forest and arable soils were observed in the top 5-cm soil layer. We propose it to be defined as a “soil stress-sensitive zone”. A strong relationship between WEOC and TOC has been reported in the literature (Zsolnay, Görlitz, 1994; Corvasce et al., 2006). As can be seen from Figure 3, the contents of CWEOC and TOC are likewise related in studied 0–50 cm soil profile. It arouses interest that a stronger correlation between these

parameters was found in arable (0.81) compared to that in forest soil (0.74).

Weigel et al. (2011) reported a strong correlation between HWEOC and TOC. They considered it to confirm that the HWEOC fraction was more related to labile carbon and thus reflected carbon changes as affected by land management. In our studies (Fig. 4) we also found a very strong correlation between these fractions: 0.98 in forest and 0.91 in arable soils. This may be explained by high content of non-humified organic material in forest soil such as lignin, lignocellulose and other carbohydrates that may be partly extracted with hot water (Wang, Wang, 2007). Another possible explanation may be the higher content of microbial, especially fungal, biomass in forest soils that is also extracted by hot water (Boyer, Groffman, 1996; Bu et al., 2011). The discovered quantitative changes of CWEOC and HWEOC fractions as well as their correlations with total



Note. R^2 – coefficient of determination; both regression lines are statistically significant at the $P < 0.01$ level; $n = 50$.

Figure 4. The dependence of hot (80°C) water extractable (HWEOC) and total organic carbon (TOC) contents in forest (a) and arable (b) soils

organic carbon content in studied soils proved the deep chemical transformation of SOM during deforestation and agricultural exploitation.

The results, as shown in Table 3, indicate that the percentage of water extractable organic carbon (CWEOC and HWEOC) in TOC was significantly lower in arable

than in forest soil (Table 3). In 0–50 cm soil layer the average percentage of CWEOC was 2.72% in forest and 2.48% in arable soil. For HWEOC fraction it was 14.08% and 9.65% in forest and arable soil, respectively. It is therefore likely that HWEOC is a much more informative indicator of soil organic matter quality than CWEOC.

Table 3. Percentage of water extractable organic carbon (WEOC, CWEOC, HWEOC) in total organic carbon (TOC)

Depth cm	Percentage (%) of CWEOC in TOC		Percentage (%) of HWEOC in TOC		Percentage (%) of WEOC in TOC	
	forest soil	arable soil	forest soil	arable soil	forest soil	arable soil
0–5	1.49	2.62	14.98	12.06	16.47	14.68
5–10	2.30	2.43	16.13	13.07	18.44	15.50
10–15	3.13	2.60	15.13	12.43	18.26	15.04
15–20	3.35	1.35	8.61	11.78	11.96	13.13
20–25	3.11	1.84	15.42	11.06	18.53	12.90
25–30	2.41	3.18	11.25	11.14	13.66	14.33
30–35	2.15	3.31	12.50	9.50	14.65	12.82
35–40	3.90	3.21	16.48	7.69	20.38	10.90
40–45	2.94	2.23	15.09	4.46	18.03	6.69
45–50	2.37	1.99	15.20	3.32	17.57	5.31

n = 5 for each soil layer

Mineralization of soil organic matter and especially its labile fractions causes the input of particles in ionic form such as H^+ , NO_3^- , NH_4^+ , HCO_3^- , HPO_4^{2-} , $H_2PO_4^-$, etc. in soil (Bedernichek et al., 2009). Taking into consideration significant change of electrical conductivity in arable compared to forest soil (Tables 1–2), it could be hypothesized that not only labile organic matter content, but the whole mineralization potential of soil has decreased. The most labile fraction of SOM – water extractable organic matter – consists of the substances with low activation energy barrier, so a depletion of WEOC pool might change the temperature sensitivity of soil (Wagai et al., 2013). Therefore the investigation of carbon dioxide emission rates from studied forest and arable soils is an important issue for future research.

Conclusions

This study was undertaken to evaluate the impact of deforestation and agriculture exploitation of *Gleyic Albeluvisol* (ABg) on distribution of total, cold and hot water extractable organic carbon in soil profile. The following conclusions can be drawn from the present research:

1. In response to forest conversion, the most prominent changes of soil chemical properties, total organic carbon (TOC), cold water extractable organic carbon (CWEOC) and hot water extractable organic carbon (HWEOC) contents were observed in the top 5-cm soil layer. We proposed it to be defined as soil stress-sensitive zone.

2. Over 70 years of agricultural exploitation of the deforested soil caused dramatic changes in its properties. Well stratified organoprofile of forest soil was simplified. The relationship between TOC content and soil depth in arable soil can be approximated not by the exponential equation as in forest, but by a simple linear function ($r = -0.97$, $P < 0.001$; $n = 10$).

3. In studied 0–50 cm layer of arable soil TOC content decreased by 32%, CWEOC by 23% and

HWEOC by 74% compared to forest soil ($n = 50$). That confirmed a high informative role of HWEOC fraction.

4. The contents of CWEOC and TOC in studied 0–50 cm soil profile were closely related. There was a stronger correlation between these parameters in arable soil ($r = 0.81$, $P < 0.01$) than in forest soil ($r = 0.74$, $P < 0.01$). Additionally, very strong correlations between HWEOC and TOC contents were observed: 0.91 ($P < 0.01$) in arable and 0.98 ($P < 0.01$) in forest soils.

5. Electrical conductivity of soil:water suspension significantly decreased in arable soil in comparison to forest soil: 34 *versus* 75 $\mu\text{Sm cm}^{-1}$. This fact might indicate the decrease of the whole mineralization potential of soil under the influence of deforestation and agriculture exploitation.

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Suminė šaltu ir karštu vandeniu ekstrahuota organinė anglis dirvožemio profilyje: žemėnaudos pakeitimo įtaka

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Santrauka

Labilios, ypač vandeniu ekstrahuotos organinės anglies kiekis yra informatyvus dirvožemio organinės medžiagos kokybės rodiklis. Tyrimo uždaviniai: 1) ištirti šaltu ir karštu vandeniu ekstrahuotos organinės anglies profilio pokyčius miško ir ariamoje žemėje, 2) nustatyti koreliaciją tarp dirvožemio organinės medžiagos labilių frakcijų ir suminės organinės anglies kiekio. Bandymai vykdyti Vakarų Ukrainoje, viršutinėje Dniestro baseino dalyje, glėjiškame balkšvažemyje (Jlg). Dirvožemio ėminiai buvo paimti iš 50 cm gylio profilio kas 5 cm. Nustatyta suminės organinės anglies, šaltu ir karštu vandeniu ekstrahuotos organinės anglies kiekiai dirvožemyje, taip pat pH (H₂O) ir dirvožemio vandens suspensijų elektrinis laidumas. Tyrimo rezultatai parodė, kad dirvožemio 0–50 cm sluoksnyje suminės organinės anglies kiekis sumažėjo 32 %, šaltu vandeniu ekstrahuotos organinės anglies – 23 %, karštu vandeniu ekstrahuotos organinės anglies – 74 %, palyginus su miško dirvožemiu, ir tai patvirtino didelę karštu vandeniu ekstrahuotos organinės anglies frakcijos įtaką. Išanalizuoti procentiniai vandeniu ekstrahuotos organinės anglies pokyčiai dirvožemio profilyje. Jie parodė, kad karštu vandeniu ekstrahuota organinė anglis yra žymiai informatyvesnis dirvožemio organinės medžiagos kokybės rodiklis, palyginus su tirpia šaltame vandenyje. Ryškiausi dirvožemio cheminių savybių – suminės, šaltame ir karštame vandenyje tirpios organinės anglies – kiekių pokyčiai, įvykę dėl miško iškirtimo, buvo nustatyti dirvožemio viršutiniame (5 cm) sluoksnyje. Šis plonas dirvožemio sluoksnis yra apibrėžiamas kaip stresui jautri zona.

Reikšminiai žodžiai: dirvožemio organinė medžiaga, labili frakcija, suminė organinė anglis, tirpi organinė anglis, vandenyje tirpi organinė medžiaga.